The role of particle-scale processes on eruption dynamics

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Simulations by Josef Dufek (now GA Tech), Helge Gonnermann (Rice)

Experiments by Ben Andrews (now Smithsonian), Bruno Cagnoli (now INGV, Rome) Marcel Staedter, Daniel Standish, Jason Wexler, Ameeta Patel, Wylie Stroberg, Rebecca Carey (UC Berkeley)

Mutiphase flows in explosive eruptions



Pyroclastic Flow

- Particulate gravity current
- Particle+Gas Flow
- Interaction with water



Challenge

- Wide range of length and time scales
- Many critical processes occur at the scale of particles and bubbles
- How to integrate the micro- and macro-scale mass, momentum and heat transfer?

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Today: 1) Review conceptual approach used in modeling and then our attempts to 2) use laboratory experiments to develop models for the micro-scale

Prior to fragmentation

- Nucleation and growth of bubbles
- Creating of permeability
- Feedback on rheology and melt composition

Numerical model

Solve equations for conservation of mass, momentum, energy at two scales



1) Conduit flow: magma (bubbles+ melt) is locally homogeneous

2) Bubble-scale:Solve for growth ofbubbles, determinerheology

Feedbacks between scales through temperature, pressure

Conduit flow

- conservation of mass, momentum, energy (include viscous dissipation; density, rheology from subgrid model)
- non-turbulent, no fragmentation,
- "single" phase magma (melt + bubbles) B
- cylindrical conduit, radial velocity is zero
- steady flow



$$-\frac{r}{2}\left(\frac{\partial p_m}{\partial z} + \rho g\right) = -\eta \frac{du_z}{dr}$$

$$\eta = \eta(\dot{\gamma}, T_m, \phi, R, c_w)$$

$$\frac{DT_m}{Dt} = \left[D_T \left(\frac{\partial^2 T_m}{\partial r^2} + \frac{1}{r} \frac{\partial T_m}{\partial r} \right) - \frac{1}{\rho_m c_{pm}} \left(\sigma_{rz} \frac{\partial u_z}{\partial r} \right) \right]$$

 $Q_{mass} = const.$

du/dr

u(r, z)

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Subgrid model: Volatile exsolution and bubble growth



Solubility of H₂0, CO₂ from Liu et al. (2005) Diffusivity of H₂0, CO₂ from Zhang and Behrens (2000)

Subgrid model: Volatile exsolution and bubble growth



Conservation of mass, momentum and energy, coupled with solubility model and modified Redlich-Kwong equation of state for water- CO_2 mixtures

$$p_g - p_m = \frac{2\gamma}{R} + 12v_R R^2 \int_R^\infty \frac{\eta_{melt}(r)}{r^4} dr.$$

Lensky et al. (2001)

$$\frac{dT_g}{dt} = \Pi \left[\rho_m c_{pm} D_T \left(\frac{\partial T_m}{\partial r} \right)_{r=R} - \sum_i \Delta H_{ev} D_i \rho_m \left(\frac{\partial c_i}{\partial r} \right)_{r=R} + \frac{R}{3} \frac{dp_g}{dt} \right] \quad \Pi = 4\pi R^2 / \left(n \ c_{pg} M_g \right)$$

$$\frac{\partial T_m}{\partial t} + v_r \frac{\partial T_m}{\partial r} = \frac{1}{r^2} \frac{\partial}{\partial r} \left(D_T \ r^2 \frac{\partial T_m}{\partial r} \right) + \frac{2 \ \eta}{\rho_m c_{pm}} \left[\left(\frac{\partial v_r}{\partial r} \right)^2 + 2 \left(\frac{v_r}{r} \right)^2 - \frac{1}{3} \left(\frac{1}{r^2} \frac{\partial}{\partial r} \left(r^2 v_r \right) \right)^2 \right]$$
Bird et al. (1960)



Basaltic eruption styles



Basaltic eruption styles



Post-fragmentation

- Collisions transfer momentum
- Collisions caused breakup and comminution
- Agglutination and sticking
- Interaction with boundaries
- Energy exchange

Multiple levels of coupling between discrete and continuous phases

Multiple levels of coupling between discrete and continuous phases



Prolonged Frictional Contact



Multiple levels of coupling between discrete and continuous phases



Instantaneous Collisions



Multiple levels of coupling between discrete and continuous phases St= τ_p / τ_f 0

Mean field multifluid equations

Continuity

$$\frac{\partial}{\partial t} ({}^{1}\alpha^{1}\rho) + \frac{\partial}{\partial x} ({}^{1}\alpha^{1}\rho^{1}u_{i}) = 0$$
Momentum

$$\frac{\partial ({}^{1}\alpha^{1}\rho^{1}u_{i})}{\partial t} + \frac{\partial ({}^{1}\alpha^{1}\rho^{1}u_{i}{}^{1}u_{j})}{\partial x_{i}} = \left[\frac{N(\alpha, e)}{{}^{\mathbf{p}}\mathbf{M}_{0}^{2}}\right] \frac{\partial ({}^{1}P)}{\partial x_{i}} + \left[\frac{1}{\mathbf{Re}}\right] \frac{\partial}{\partial x_{i}} [{}^{1}\tau_{ij}] + \left[\frac{1}{\mathbf{St}}\right] ({}^{1}u_{i} - {}^{2}u_{i}) + \left[\frac{1}{\mathbf{Fr}_{d}^{2}}\right] \alpha \hat{e}_{g}$$
Thermal Energy

$${}^{1}\rho^{1}\mathbf{c}_{p} \left[\frac{\partial^{1}\mathbf{T}}{\partial \mathbf{t}} + {}^{1}\mathbf{U}_{i}\frac{\partial^{1}\mathbf{T}}{\partial x_{i}}\right] = \left[\frac{1}{\mathbf{Pe}}\right] \frac{\partial^{1}\mathbf{q}}{\partial^{1}x_{i}} + \left[\frac{1}{\mathbf{Th}}\frac{1}{\mathbf{Th}}\right] ({}^{2}\mathbf{T} - {}^{1}\mathbf{T})$$

Details of constitutive models, equations of state, turbulence models, in Dufek and Bergantz, *Theoretical and Computational Fluid Dynamics* (2007)

Mean field multifluid equations

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Numerical implementation adapted from the MFIX (**M**ultiphase Flow with Interface eXchange) code developed by US Dept Energy.

Sub-grid scale thermo-mechanical processes

1) Heat transfer from particles to water

2) Mass and momentum transfer when particles collide with each other

1. Interaction with water

Hot flows, when they enter water, generate steam

- How much?
- How fast?



• Effects of steam generation?

Measurement of steam production rate



 Measure mass of stream released
 Measure time clasts float (results in Dufek, Manga Staedter, *J Geophys Res* 2007)

Measurement of steam production rate



Continuity

 $\frac{\partial}{\partial t}({}^{w}\alpha{}^{w}\rho) + \frac{\partial}{\partial x_{i}}({}^{w}\alpha{}^{w}\rho{}^{w}U_{i}) = \underbrace{-R_{v}}_{\text{Mass loss due to}}$ $\frac{\partial}{\partial t}({}^{g}\alpha{}^{g}\rho) + \frac{\partial}{\partial x_{i}}({}^{g}\alpha{}^{g}\rho{}^{g}U_{i}) = \underbrace{+R_{v}}_{\text{Mass gain due to}}$ $\frac{\partial}{\partial t}({}^{p}\alpha{}^{p}\rho) + \frac{\partial}{\partial x_{i}}({}^{p}\alpha{}^{p}\rho{}^{p}U_{i}) = 0$

Multiphase equations

Momentum

$$\frac{\partial}{\partial t}({}^{g}\alpha {}^{g}\rho U_{i}) + \frac{\partial}{\partial x_{i}}({}^{g}\alpha {}^{g}\rho {}^{g}U_{i} {}^{g}U_{j}) = \frac{\partial {}^{g}P}{\partial x_{i}}\delta_{ij} + \frac{\partial {}^{g}\tau_{ij}}{\partial x_{j}} + {}^{g}I_{i} + {}^{g}\alpha {}^{g}\rho g_{i} + \underbrace{R_{v} {}^{g}U_{i}}_{\text{Momentum gain to}} \\ \frac{\partial}{\partial t}({}^{w}\alpha {}^{w}\rho U_{i}) + \frac{\partial}{\partial x_{i}}({}^{w}\alpha {}^{w}\rho {}^{w}U_{i} {}^{w}U_{j}) = \frac{\partial {}^{w}P}{\partial x_{i}}\delta_{ij} + \frac{\partial {}^{w}\tau_{ij}}{\partial x_{j}} + {}^{w}I_{i} + {}^{w}\alpha {}^{w}\rho g_{i} - \underbrace{R_{v} {}^{g}U_{i}}_{\text{Momentum gain to}} \\ - \underbrace{R_{v} {}^{w}U_{i}}_{\text{Momentum loss due}} \\ + \underbrace{R_{v} {}^{g}U_{i}}_{\text{Momentum loss due to phase change}}$$

$$\frac{\partial}{\partial t}({}^{p}\alpha{}^{p}\rho U_{i}) + \frac{\partial}{\partial x_{i}}({}^{p}\alpha{}^{p}\rho{}^{p}U_{i}{}^{p}U_{j}) = \frac{\partial{}^{p}P}{\partial x_{i}}\delta_{ij} + \frac{\partial{}^{i}\tau_{ij}}{\partial x_{j}} + {}^{p}I_{i} + {}^{p}\alpha{}^{p}\rho g_{i}$$

Thermal Energy

$${}^{w}\alpha {}^{w}\rho {}^{w}c_{p}\left(\frac{\partial {}^{w}T}{\partial t} + {}^{w}U_{i}\frac{\partial {}^{w}T}{\partial x_{i}}\right) = \frac{\partial {}^{w}q}{\partial x_{i}} + \overline{H}_{wg} - \underbrace{\overline{H}_{wp}}_{\text{Mean interphase}} - \underbrace{\overline{H}_{wp}}_{\text{Mean interphase}} + \underbrace{\overline{S}}_{\text{heat transfer}} + \underbrace{\overline{S}}_{\text{heat interphase}} + \underbrace{\overline{S}}_{\text{heat of}} + \underbrace{\overline{S}}_{\text{vaporization}} +$$



Water (100 m deep)

grid 2 m x 10 m; time step , 0.1 s initial velocity 50 m/s initial concentration 0.1 initial sizes: 50% is 1 cm, 50% is 0.1 mm temperature 650 C







Volume Fraction of Particles with Subgrid Steam Production







Edmonds and Herd, Geology (2005)

- 0.6% flow forms landward-directed base surge (Edmonds et al. (2006) estimate a volume of 0.75%)
- Landward directed flow is dry

2. Studying particle collisions



Measure velocity before and after collision; whether particle bounces
Variables: angle φ, velocity, mass, substrate

Example (water substrate)



• Extract quantitative information . . .

Pumice clasts can break if collisions are energetic enough



Will large pumice clasts breakup before exiting volcanic conduits?

Analytical model



- Assumptions: choked flow (exit velocity is the speed of sound in a dusty gas)
- Dissipation of granular energy balanced by production owing to shear



Lagrangian particles



Detailed expressions in Dufek, Wexler and Manga, J Geophys Res (2009)

Lagrangian analysis



Most large clasts are disrupted for fragmentation > few 100 m



Conclusions

- Steam generation at the scale of particles change flow behavior qualitatively
- Large pumice suggests shallow fragmentation





Conclusions

- Integration models at different scales has a long history (e.g., turbulence) – and necessary for many volcanic processes because of the vast range of length and time scales that matter
- Experimental measurements can be used to link the micro- and macro-scale
- The "new" thermo-mechanical processes and properties (ash production, vaporization of water, boundary conditions) we have included matter in simulations - qualitatively and quantitatively